

# The impact of clouds on radical concentrations: Observations of OH and HO<sub>2</sub> during HCCT2010

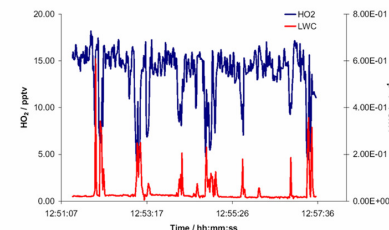
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## 1. Rationale

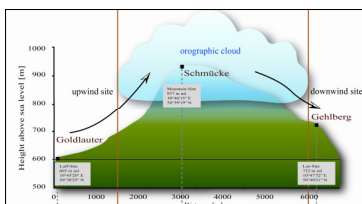
- Clouds occupy ~ 15% of the troposphere.
- Modelling studies have shown that clouds can influence gas-phase chemistry either directly through efficient uptake into cloud droplets or indirectly through the separation of precursors between phases (Lelieveld & Crutzen, 1990, 1991; Tilgner et al., 2005). The reduction of gas-phase oxidants can decrease the self-cleansing capacity of the troposphere and change the tropospheric lifetimes of VOCs.
- Only very few experimental data exist to support model results (Mauldin et al, 1997; Frost et al, 1999; Commane et al, 2010).
- ~ 75% reduction in [HO<sub>2</sub>] in cloud during aircraft flights has been observed previously (Commane et al, 2010) – see figure 1.



**Figure 1**, time-series of [HO<sub>2</sub>] and liquid water content during an aircraft flight over Africa, taken from Commane et al., (2010).

## 2. HCCT campaign methodology

The Hill Cap Cloud Thuringer – 2010 (HCCT-2010) campaign aimed to further investigate the connection of HO<sub>x</sub> reduction in clouds and microphysical parameters such as liquid water content.



**Figure 2**, Schematic of the HCCT-2010 campaign set-up that took place in Thuringer Wald, Germany. Experiment consisted of 3 sites: Goldlauter (upwind), Schmücke (summit), and Gehlberg (downwind) and aimed to study cloud processing of gas-phase species and aerosols.



**Figure 3**, photo of FAGE instrument at the Schmücke site during a cloud event. Radicals were sampled from the top of the 22 m tower alongside ancillary measurements.

Laser induced fluorescence (LIF) at low pressures, known as the FAGE technique, was used to detect the gas-phase OH and HO<sub>2</sub> radical concentrations at the summit site both in-cloud and out of cloud. The data were compared to simple model predictions.

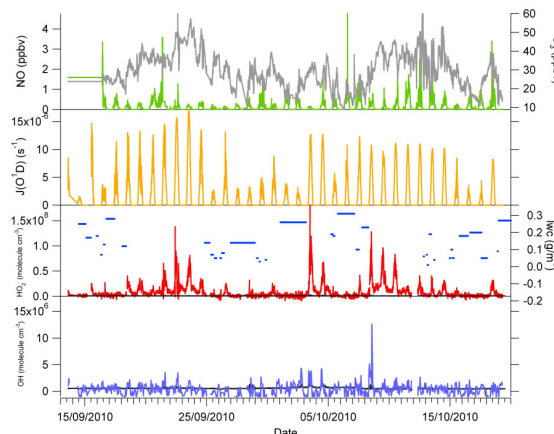
## 6. References

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## 7. Acknowledgements

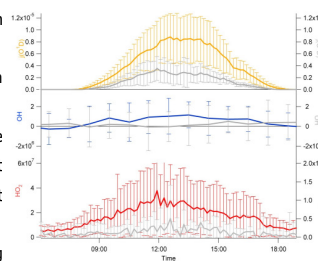
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## 3. Results



**Figure 4**, time-series of [OH], LOD<sub>OH</sub>, [HO<sub>2</sub>], lwc, J(O<sup>1</sup>D), [NO] and [O<sub>2</sub>] at the Schmücke site from 14<sup>th</sup> September – 19<sup>th</sup> October 2010.

- Many cloud events occurring both during the day and night.
- ~ 65% reduction in J(O<sup>1</sup>D) on average during cloud events.
- Low concentrations of OH were observed out of cloud, but concentrations were below the limit of detection in-cloud.
- ~ 90% reduction in [HO<sub>2</sub>] during cloud events observed on average.



**Figure 5**, average diurnal of J(O<sup>1</sup>D), [OH] and [HO<sub>2</sub>] during cloud events (grey) and out of cloud (colour).

## 4. Simple model predictions of radicals

[OH] may be estimated by balancing the rate of production with the rate of loss:

$$[\text{OH}]_{\text{loss}} = \frac{p(\text{OH}) + k[\text{HO}_2][\text{NO}] + k[\text{HO}_2][\text{O}_3]}{k[\text{CO}] + k[\text{CH}_4]}$$

where p(OH) is the primary production of OH from the photolysis of O<sub>3</sub> and subsequent reaction of O(<sup>1</sup>D) with H<sub>2</sub>O (v).

Quadratic expression used to estimate [HO<sub>2</sub>]:

$$0 = [\text{HO}_2]^2 (2k_{\text{HO}_2+\text{HO}_2}) + [\text{HO}_2] (k_{\text{HO}_2+\text{NO}}[\text{NO}] + k_{\text{HO}_2+\text{RO}_2}[\text{RO}_2] + k_{\text{HO}_2+\text{O}_3}[\text{O}_3]) - (k_{\text{CO/OH}}[\text{CO}][\text{OH}] + 2j(\text{HCHO})[\text{HCHO}] + k_{\text{RO}_2+\text{O}_2})$$

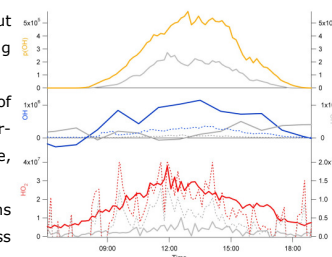
$$0 = a[\text{HO}_2]^2 + b[\text{HO}_2] - c$$

$$[\text{HO}_2] = \frac{-b \pm \sqrt{b^2 - 4ac}}{2a}$$

- Lower concentrations of [OH] predicted out of cloud than observed suggesting missing OH sources.
- Model is able to replicate [HO<sub>2</sub>] out of cloud (red dashed line, Fig. 6) but over-estimates [HO<sub>2</sub>] in-cloud (grey dashed line, Fig. 6)
- To replicate [HO<sub>2</sub>] in-cloud observations (solid grey line, Fig. 6), a first order loss rate of HO<sub>2</sub> to clouds = 0.2 s<sup>-1</sup> is needed, suggesting a reactive uptake coefficient γ = 0.0125 (at cloud droplet surface area = 1.5 × 10<sup>-3</sup> cm<sup>2</sup>cm<sup>-3</sup>), calculated from:

$$\gamma = \frac{4k}{\omega A}$$

Where k is the first order loss rate (s<sup>-1</sup>), ω is the mean molecular speed of HO<sub>2</sub> (cm s<sup>-1</sup>), and A is the surface area (cm<sup>2</sup>cm<sup>-3</sup>).



**Figure 6**, average diurnal of p(OH), [OH] and [HO<sub>2</sub>] during cloud events (grey) and out of cloud (colour). Model estimates (without heterogeneous loss to cloud droplets), in cloud (grey dash) and out of cloud (colour dash) for OH and HO<sub>2</sub> are also shown.

## 5. Conclusions and future work

- HO<sub>2</sub> concentrations were significantly depleted in cloud, with concentrations ~ 90 % lower relative to the out of cloud observations; an OH signal above the noise of the instrument was not observed during cloud events. These results suggest that heterogeneous processes in clouds do perturb the gas-phase radical chemistry.
- Further investigations studying the uptake of HO<sub>2</sub> during different cloud events.
- Does the uptake depend upon cloud microphysical properties?
- Comparison of observations with a multiphase chemistry model – Master Chemical Mechanism (MCM) coupled with Chemical Aqueous Phase Radical Mechanism (CAPRAM).